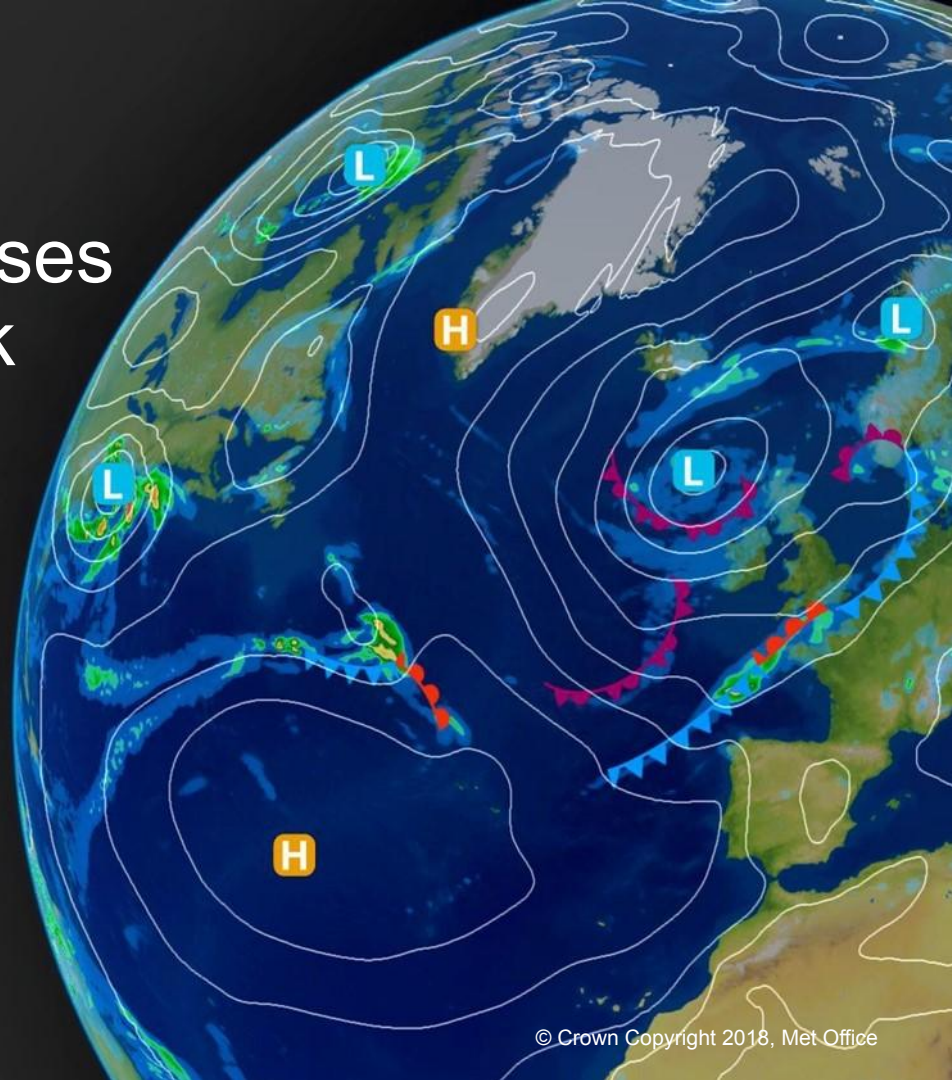


Understanding sea-ice biases in the seamless framework

By Dan Copsey and Tim Graham



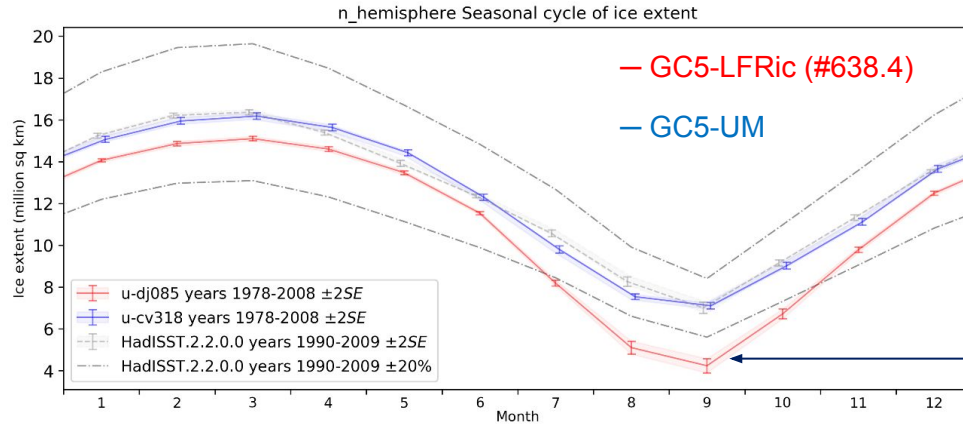
The new GC5-LFRic model

(compared to the GC5-UM model)



- Uses a new cube-sphere grid.(GC5-UM is regular lat-long)
- Includes a new dynamical core: Gung-Ho. GC5-UM uses EndGame.
- Maintains almost all the same physics parameterisations as GC5-UM (so should perform the same).
- Is still coupled to the same ocean and sea ice models as GC5-UM (GOSI9).

The GC5-LFRic sea ice issue

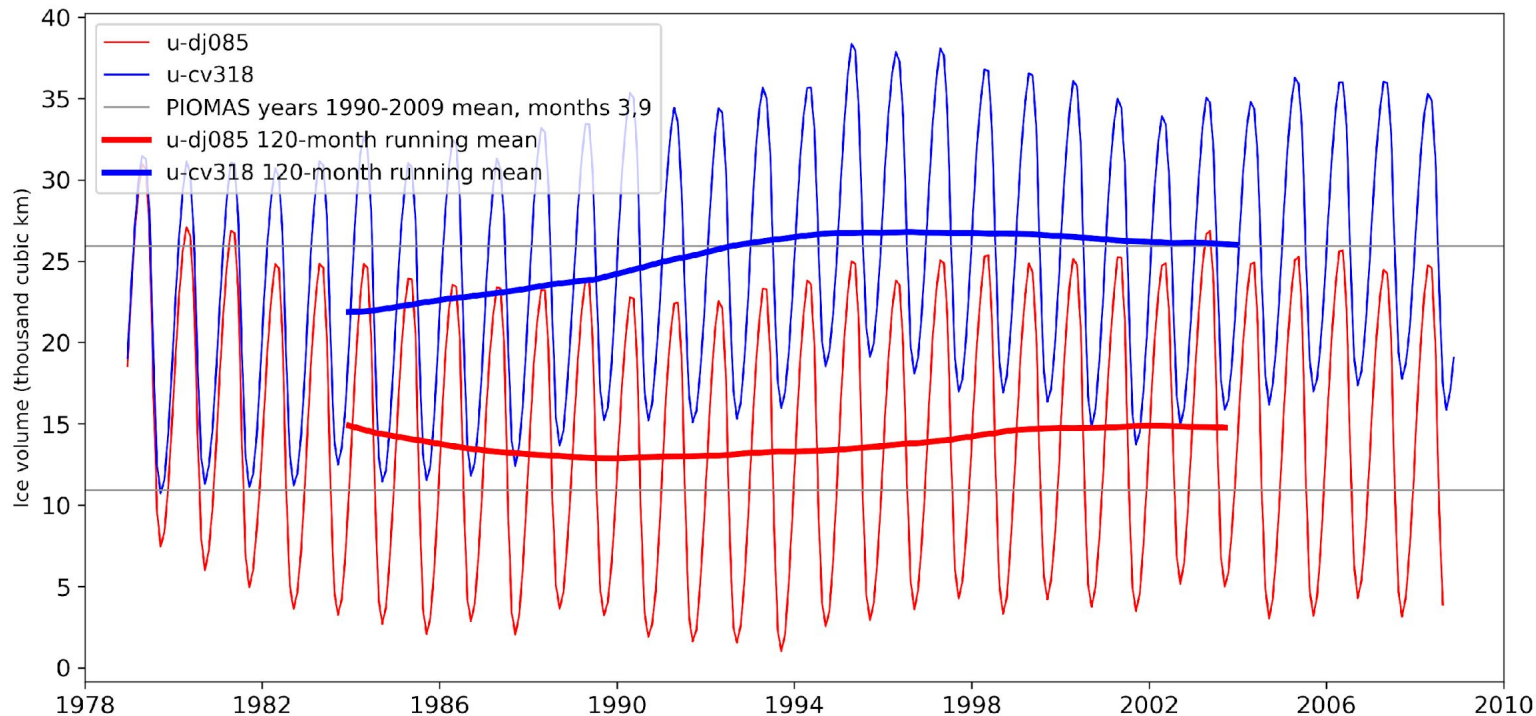


Model = C144 (or N216) ORCA025
30 year means (present day)

The new GC5-LFRic model experiences excessive Arctic sea ice loss during the summer.

— GC5-LFRic (#638.4)
— GC5-UM

Time series of sea ice volume (Northern Hemisphere)



Sea ice volume substantially reduced throughout the year for GC5-LFRic, particularly in Northern Hemisphere.

Slide from Emma Fiedler

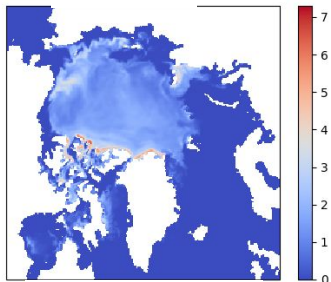
Identifying similar sea ice loss in coupled NWP simulations

Model = C224 (or N320) ORCA025
6 day means
(initialised on 30th June 2021)

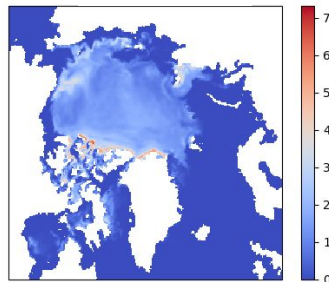
Change in sea ice volume after 6 days

Sea ice volume (sivolu) 20210630

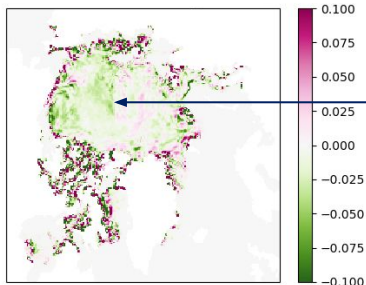
GC5-UM



GC5-LFRic



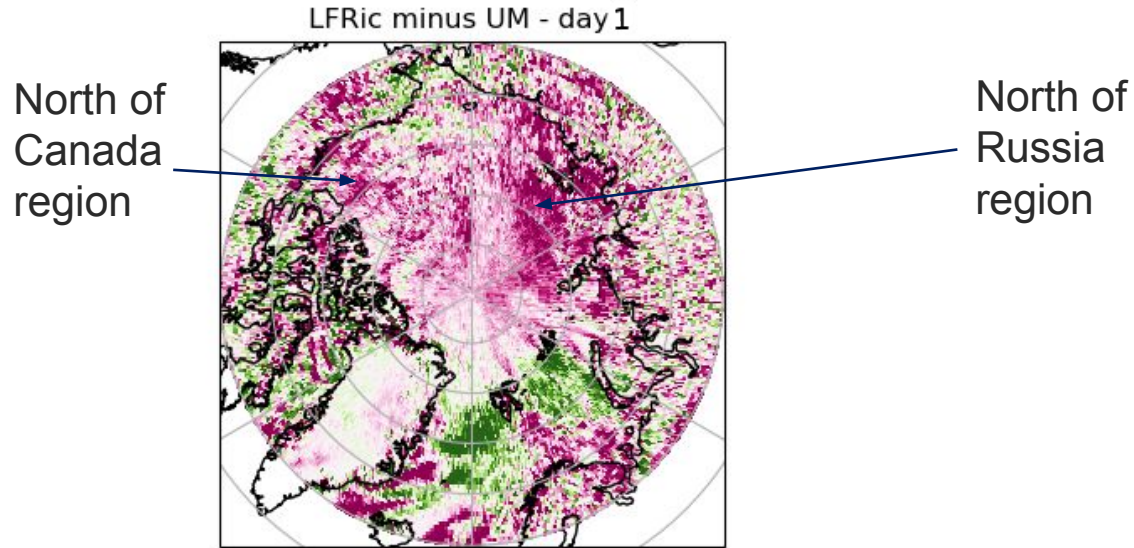
LFRic minus UM



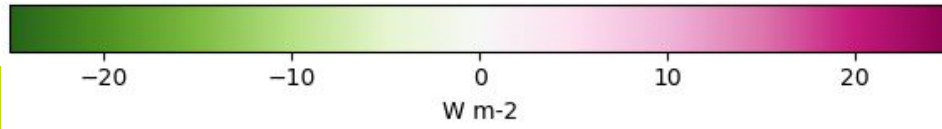
The new GC5-LFRic model
loses sea ice during the 6 days
of the simulation.

Analysing net downward surface shortwave fluxes

- Downward net surface shortwave



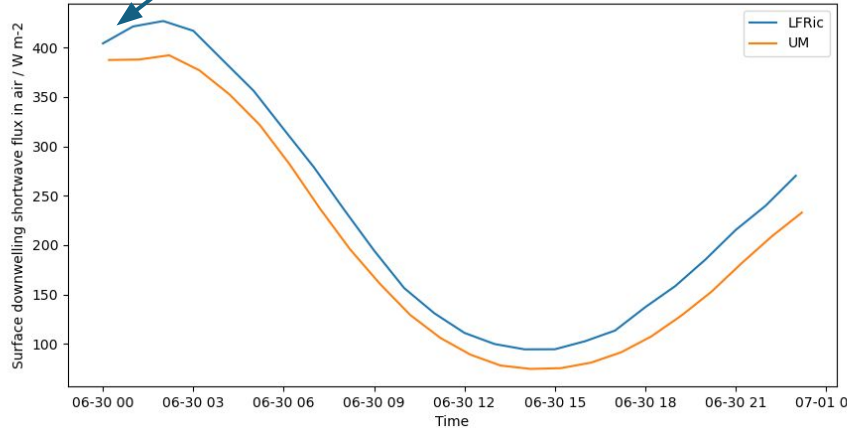
LFRic has increased downward shortwave leading to more sea ice melt in these regions



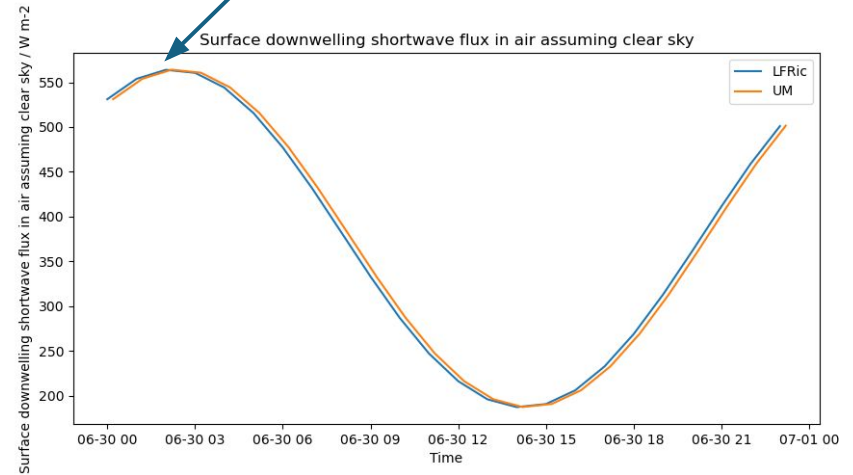
Timeseries of surface SW fluxes in region North of Russia over first day

LFRic has an immediate increase in surface shortwave at the first call to radiation before any cloud changes can have an effect

Surface downwelling shortwave flux in air



Clear sky is virtually identical suggesting the problem is with the clouds (or how the clouds are represented by radiation)



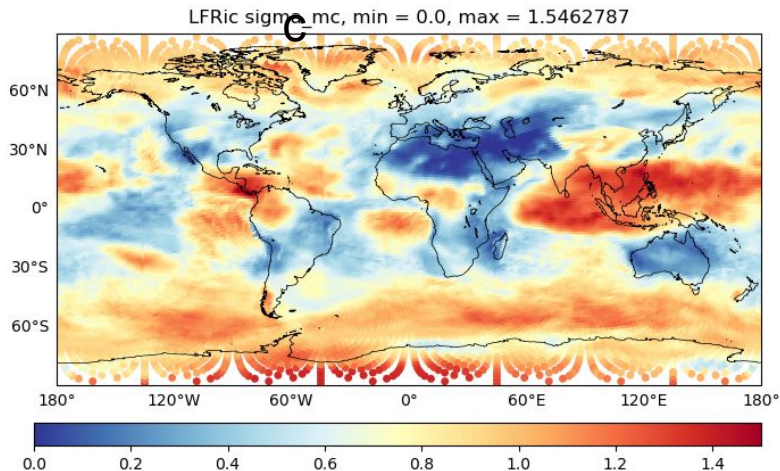
Two theories

- The cloud inhomogeneity is different between UM and LFRic
- The cloud droplet effective radius is different between UM and LFRic

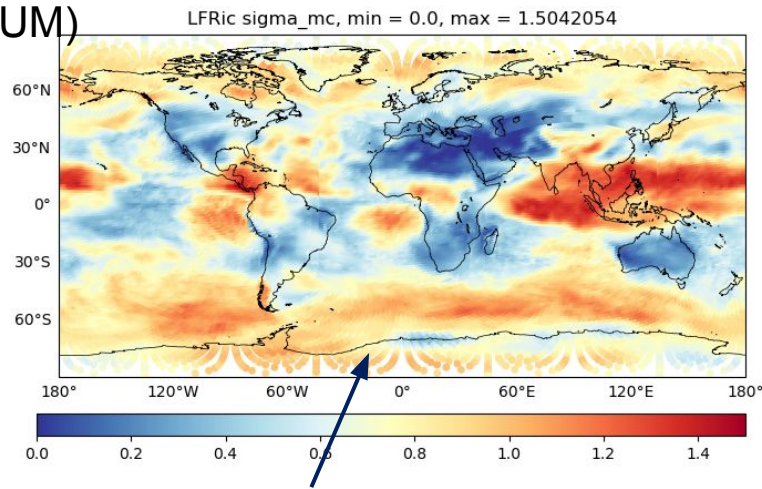
How does the cloud inhomogeneity parameter change (cloud Fractional Standard Deviation - FSD)

The FSD is a measure of how fragmented the clouds are. Larger = more fragmented = more shortwave reaches surface.

GC5-LFRi



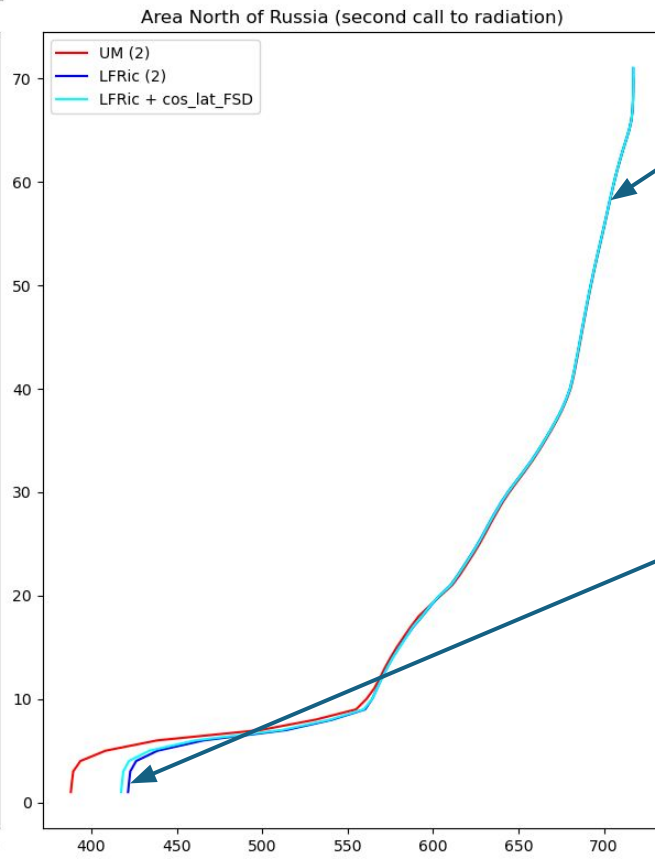
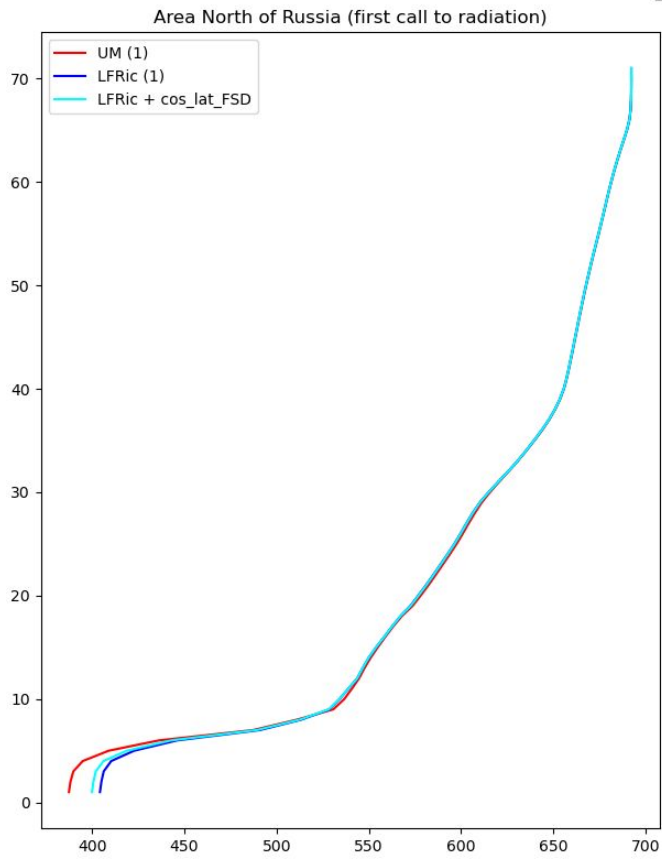
GC5-LFRic but with latitudinal FSD (like the UM)



Adding latitudes to FSD calculations reduce FSD near poles. Clouds are more uniform.

How does changing FSD change the solar flux as it descends through the atmosphere

downward_shortwave



Cyan line = Using latitudinal dependent FSD

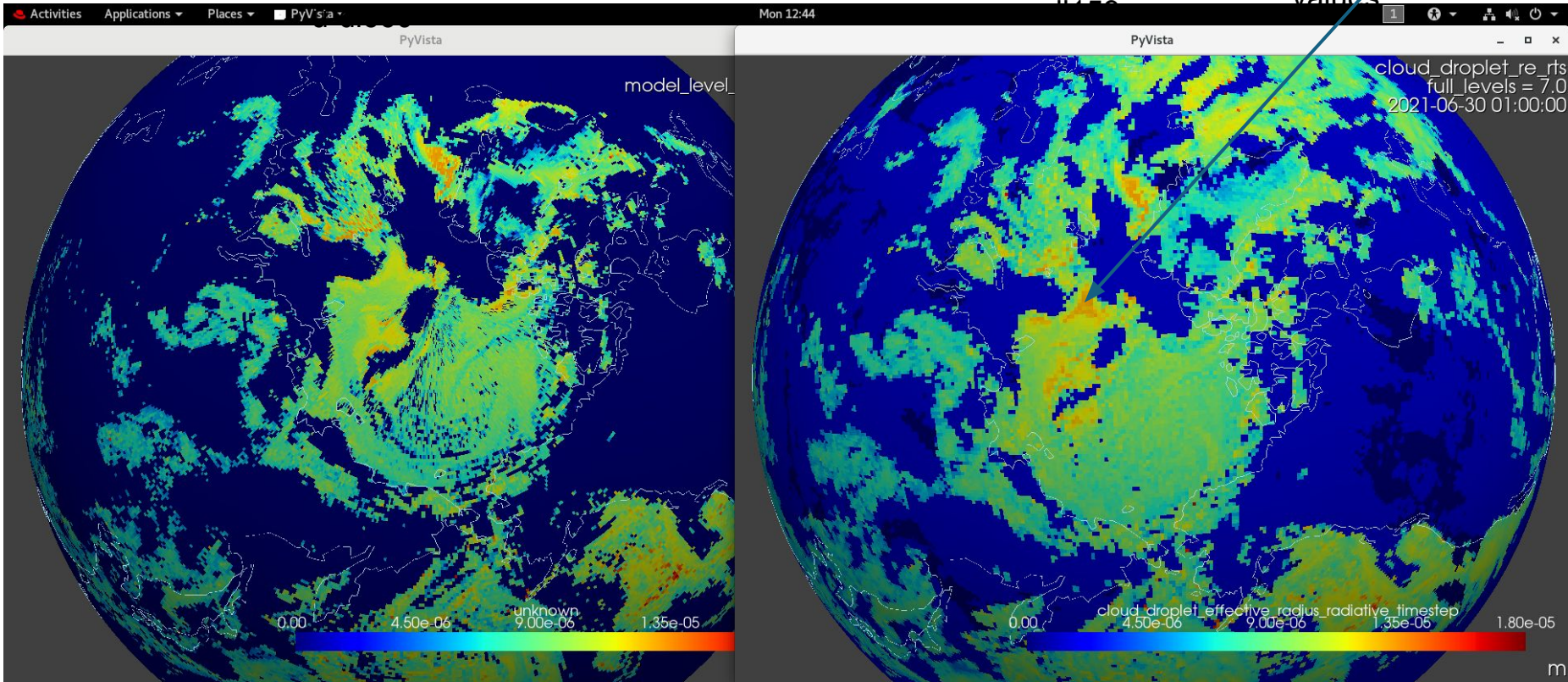
Changing cloud inhomogeneity can increase downward shortwave (but only by a little).

Cloud droplet effective radius on level 7 at first call to radiation

UM

LFRic

LFRic has larger values

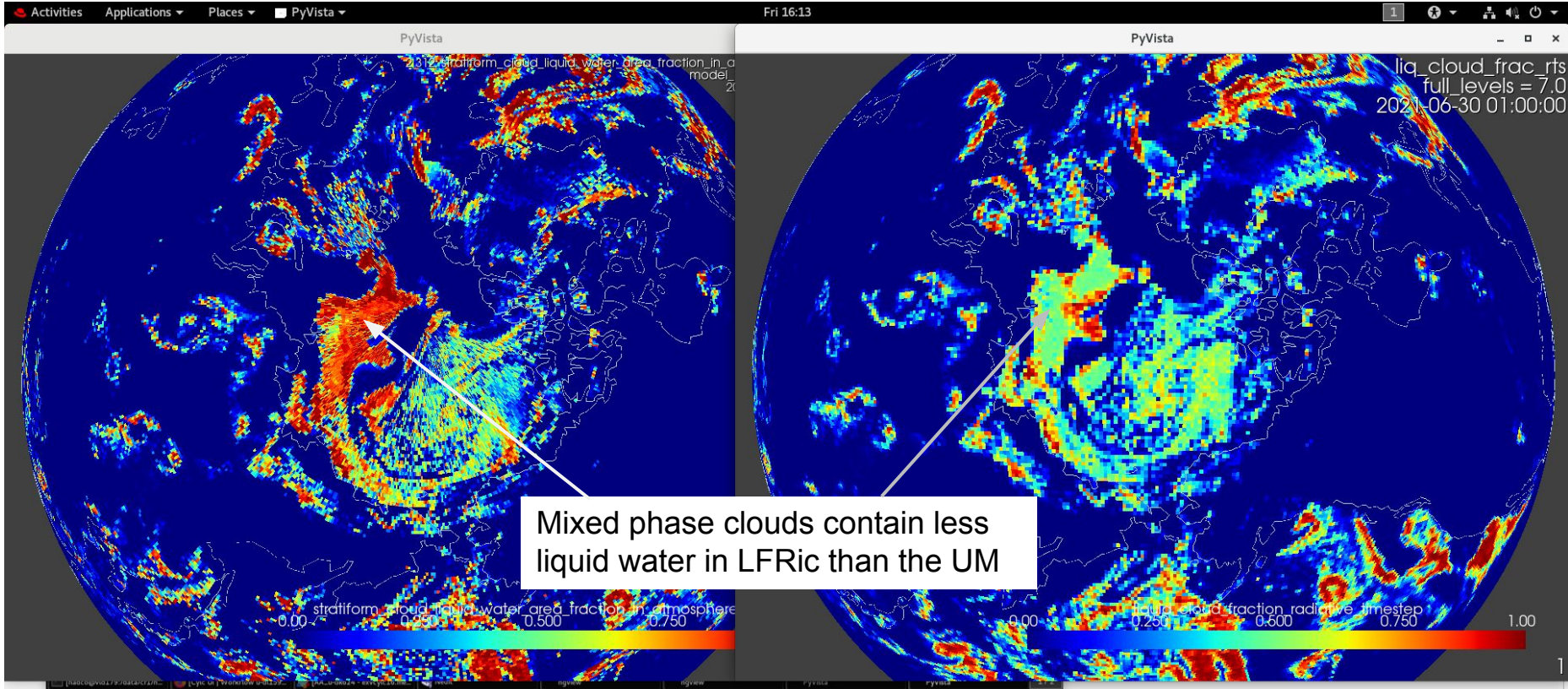


Liquid cloud fraction (as used by radiation) on level 7 at first call to radiation

UM

LFRic

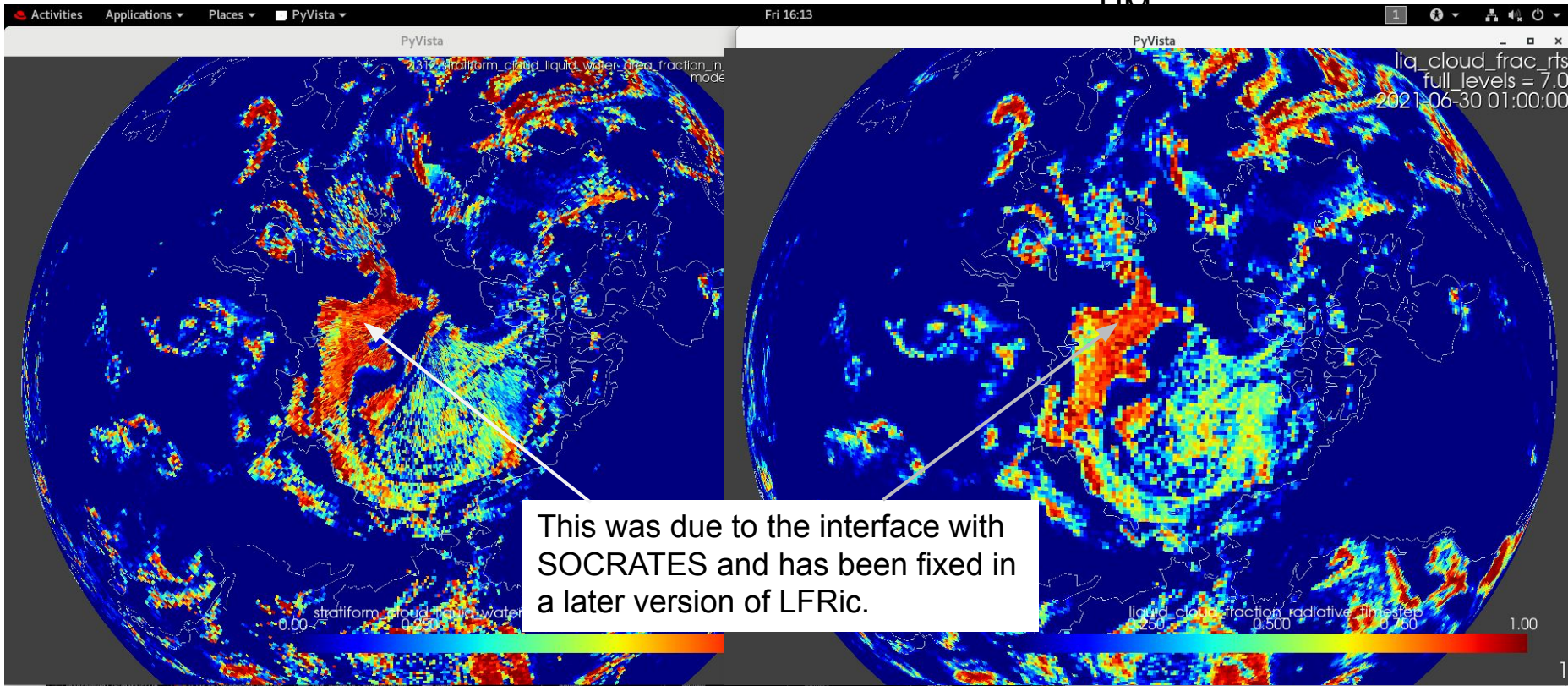
LFRic has less cloud



Fixed liquid cloud fraction (as used by radiation) on level 7 at first call to radiation

UM

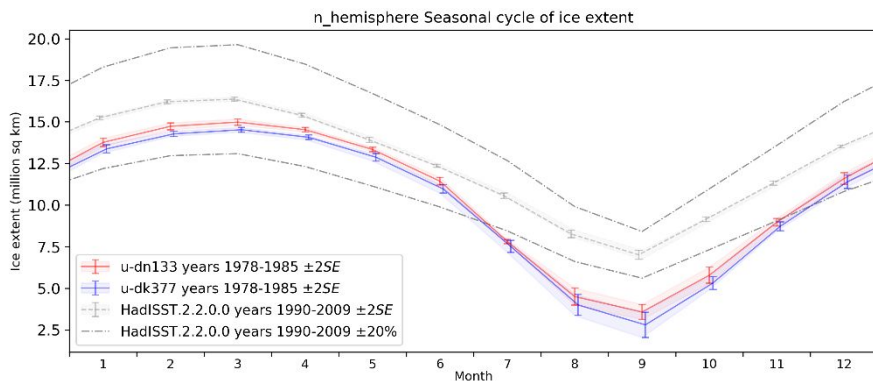
LFRic u-dn080 LFRic now matches



Summary of coupled NWP investigations

- Over the Arctic, in locations where there are mixed phase clouds, LFRic underestimated the contribution from liquid cloud to the radiation scheme. This has now been fixed because of this investigation.
- Over the Arctic the cloud inhomogeneity scheme behaves differently because it is dependent on the grid spacing and that is different between the UM and LFRic. This investigation has shown that some cloud inhomogeneity tuning can go some way to improve the model.

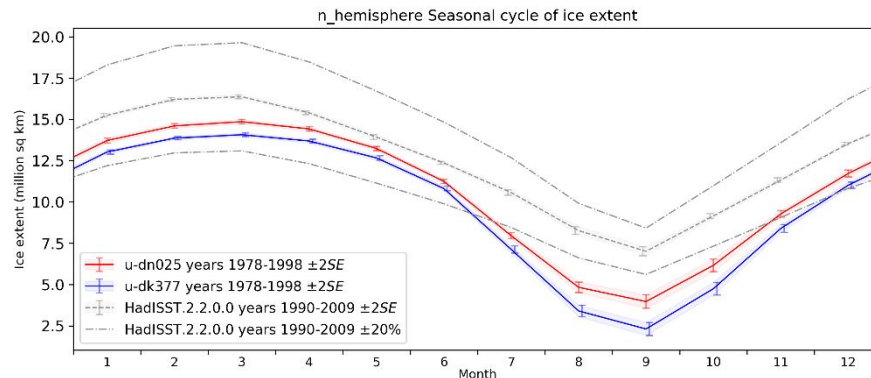
The effect of the mixed phase fix



Blue = GC5-LFRic control

Red = GC5-LFRic with mixed_phase_fix

The effect of adding latitude dependence to cloud inhomogeneity calculations



Red = GC5-LFRic with latitude dependence in cloud inhomogeneity

Coding up a new albedo scheme for melt ponds using Malinka et al.



Snell's law for refracted angle in water

$$\sin(\theta_{water}) = \frac{n_{air}}{n_{water}} \sin(\theta_{air})$$

Fresnel equations for reflection off of a liquid surface

$$R_s = \frac{n_{air} \cos(\theta_{air}) - n_{water} \cos(\theta_{water})}{n_{air} \cos(\theta_{air}) + n_{water} \cos(\theta_{water})}$$

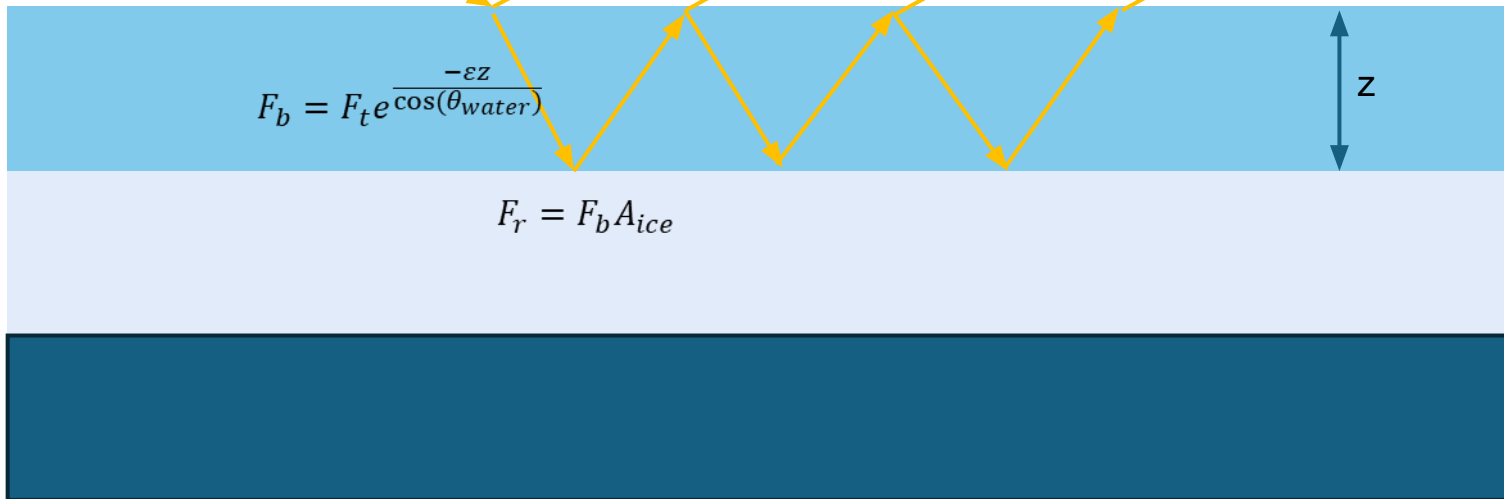
Melt pond

$$F_b = F_t e^{\frac{-\epsilon z}{\cos(\theta_{water})}}$$

Sea ice

$$F_r = F_b A_{ice}$$

Ocean



Coding up a new albedo scheme for melt ponds using Malinka et al.

- Makshtas and Podgorny initially combined the above equations from the previous slide into 1 equation and then Malinka et al rewrote them into a simple form:

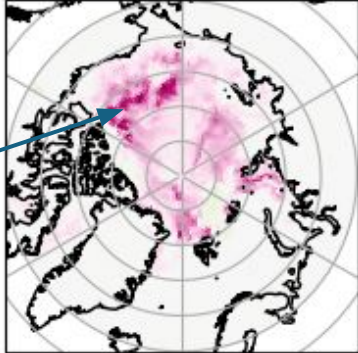
$$A(\mu_0) = R^F(\mu_0) + \frac{T^F(\mu_0) \exp(-\varepsilon_w z / \mu_0^w) f_{\text{out}}(\varepsilon_w z) A_b}{n^2 (1 - A_b f_{\text{in}}(\varepsilon_w z))}, \quad (1)$$

This does the direct albedo but for the diffuse albedo this is integrated over all incident angles. A best-fit curve is applied as a full integral would be too slow for a physics parameterisation.

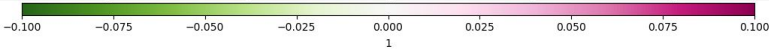
Effect of Malinka et al melt pond albedo scheme on coupled NWP case study (30/06/2021)

Surface

Diff mean over 6 days

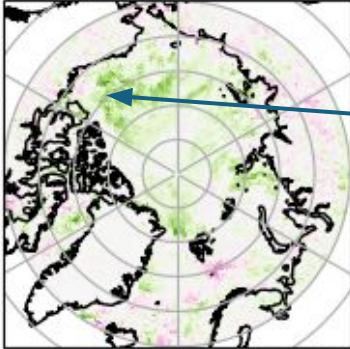


Surface albedo increases in locations of melt ponds

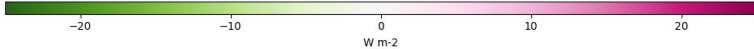


Net down surface sh

Diff mean over 6 days



Net downward shortwave decreases

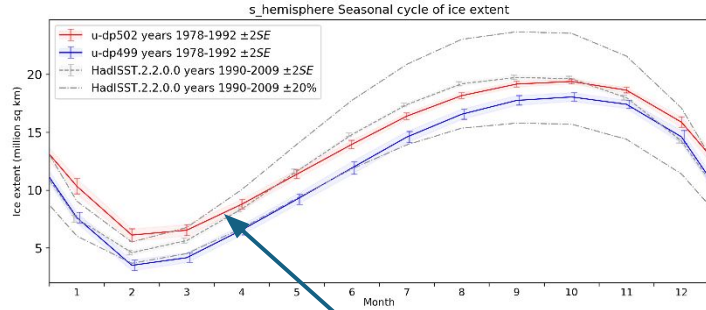
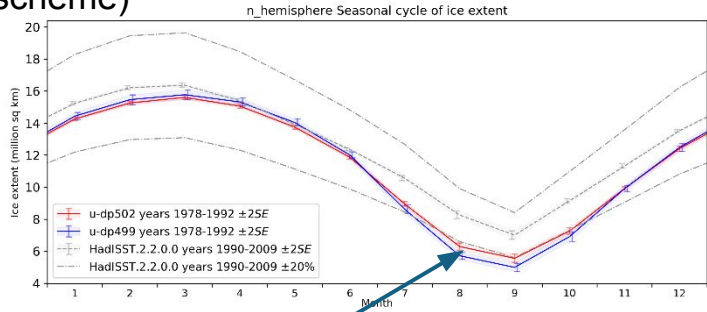


Effect of Malinka et al melt pond albedo scheme on 14 year

climate run

u-dp499 = Control (GC5-lfric with no sea ice albedo changes)

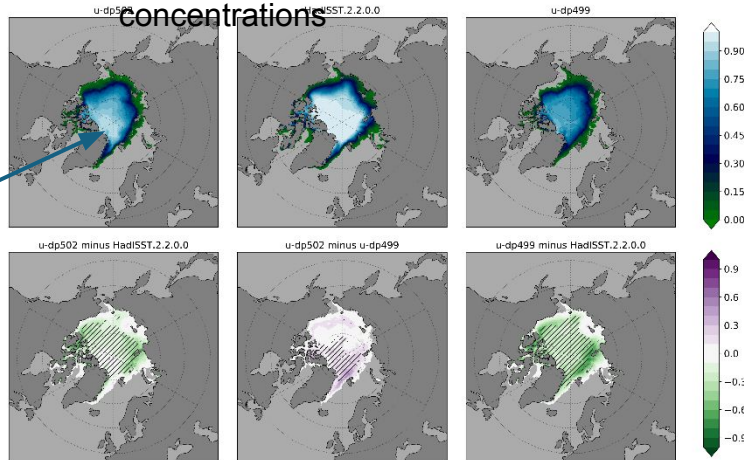
u-dp502 = Experiment (GC5-lfric with zenith angle albedo changes plus Malinka et al melt pond albedo scheme)



Increasing Northern Hemisphere sea ice a bit

The increases in Arctic sea ice are in the right place (near Greenland)

September sea ice concentrations



Increasing Southern Hemisphere sea ice a lot more (maybe too much)

Summary

- We have used coupled NWP experiments to understand biases and test fixes for problems in climate experiments.
- This massively reduced the cost of the investigations and testing allowing quicker understanding.
- The coupled NWP investigation into LFRic vs UM differences have revealed two discrepancies that make GC5-LFRic have less sea ice than GC5-UM when the same model is used for climate simulations:
 - A difference in the mixed phase clouds
 - A difference in the cloud inhomogeneity (FSD)
- A new melt pond albedo scheme has been developed (using Malinka et al) that is more physically realistic and shows improvements (increases) to the sea ice in both coupled NWP and climate simulations.
- Package tests are underway that combine all these improvements (and more) for GC6.0

Questions?